Observation and modeling of the evolution of Texas power plant plumes

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Abstract

During the second Texas Air Quality Study 2006 (TexAQS II), a full range of pollutants was measured by aircraft in eastern Texas during successive transects of power plant plumes (PPPs). A regional photochemical model is applied to simulate the physical and chemical evolution of the plumes. The observations reveal that SO_2 and NO_y were rapidly removed from PPPs on a cloudy day but not on the cloud-free days, indicating efficient aqueous processing of these compounds in clouds. The model reasonably represents observed NO_x oxidation and PAN formation in the plumes, but fails to capture the rapid loss of SO_2 (0.37 $hour^{-1}$) and NO_y (0.24 $hour^{-1}$) in some plumes on the cloudy day. Adjustments to the cloud liquid water content (QC)

and the default metal concentrations in the cloud module could explain some of the SO_2 loss. However, NO_y in the model was insensitive to QC. These findings highlight cloud processing as a major challenge to atmospheric models. Model-based estimates of ozone production efficiency (OPE) in PPPs are 20-50% lower than observation-based estimates for the cloudy day.

1 Introduction

Power plants are the leading point source emitters of SO_2 and oxides of nitrogen (NO_x=NO+NO_2) (EPA, 2009). The large amount of SO_2 and NO_x emitted from power plants has been linked to a series of environmental issues, such as acid deposition, photochemical O_3 and particulate matter (Srivastava et al., 2004; Ryerson et al., 2001; Brock et al., 2003; Flues et al., 2002). Various regulations and market-based policies have been implemented to reduce these emissions, including the Acid Rain Program (EPA, 2005) and the NO_x State Implementation Plan Call (NO_x SIP Call) (EPA, 2004) in the United States. Power plants are among the most accurately measured emission sources in the U.S. national emission inventory due to direct smoke stack measurements by Continuous Emission Monitoring Systems (CEMS). Good agreement has been found in comparing power plant emissions reported by CEMS with airborne measurements of power plant plumes (PPPs) (Frost et al., 2006) and with satellite measurements of NO_2 (Kim et al., 2006).

The emissions, transport, and chemical evolution of pollutants from power plants have been investigated by multiple observational and modeling methods (Ryerson et al., 1998; Neuman et al., 2004; Godowitch et al., 2008a; Frost et al., 2006; Kim et al., 2006; Sillman, 2000). Airborne measurement of chemical composition and meteorological parameters in PPP transects have

been conducted in several field campaigns over North America (Trainer et al., 1995; Ryerson et al., 1998; Springston et al., 2005; Neuman et al., 2009).

SO_2 freshly emitted from power plant stacks is quickly diluted and undergoes chemical evolution during plume transport. Previous aircraft measurements in PPPs have revealed that gas-phase SO_2 oxidation is the key pathway for the SO_2 removal and the particle growth in PPPs in the absence of clouds (Brock et al., 2002, 2003; Springston et al., 2005). SO_2 can also readily dissolve in cloud water and then convert to sulfate via aqueous reactions.

Several previous field studies have investigated the chemical evolution and lifetime of NO_x, ozone production efficiency, and the loss rate of reactive nitrogen in PPPs (Ryerson et al., 1998; Springston et al., 2005; Neuman et al., 2009). Some studies have reported the rapid loss of NO_y in PPPs which they primarily attributed to HNO_3 loss (Neuman et al., 2004;Nunnermacker et al., 2000), but others have not (Ryerson et al., 2003).

Power plants are significant contributors to NO_x and SO_2 emissions and high ozone concentrations in eastern Texas. A NOAA WP-3 aircraft performed successive downwind transects of PPPs in this region during several flights as part of the summer 2006 Second Texas Air Quality Study (TexAQS II) (Parrish et al., 2009). The instruments aboard the WP-3 measured a full range of trace gases, aerosol parameters, and meteorological parameters at high time resolution and spatial resolution. This study utilizes the rich data source to examine whether a 3D photochemical model with a fine spatial resolution but without subgrid plume treatment can effectively simulate the chemical and physical evolution of PPPs as they disperse and transport downwind. We focus on the evolution of sulfur, reactive nitrogen, and O_3 in the plumes.

2 Airborne measurement

TexAQS II was a comprehensive observational campaign in eastern Texas from August to October, 2006, which aimed to improve scientific understanding of the sources and atmospheric processes responsible for the formation and distribution of O_3 and aerosols in the region (Parrish et al., 2009). PPPs observed during TexAQS II originated from eastern Texas coal-fired power plants with a large range of reported NO_x and SO_2 emission rates (Table 1 and Fig. 1).

The measurements and operational characteristics of the NOAA WP-3 have been summarized elsewhere (Parrish et al., 2009). Instruments aboard the WP-3 measured numerous reactive nitrogen species (NO, NO_2, HNO_3, NO_3, N_2O_5, PAN, peroxy propionyl nitrate, methacrylol peroxy nitrate), isoprene, CO_2, CO, SO_2, HCHO, major aerosol parameters, UV-VIS actinic flux, relative humidity, and temperature (Tables A1a and A1b of Parrish et al. (2009) and the references therein). The instruments used in measuring major gas-phase species are summarized in Table S2 (in supplementary materials). The time resolution of most instruments was 1 second, equal to approximately 100m spatial resolution at typical WP-3 flying speeds.

Coal-fired power plants are major sources of SO_2 and NO_x, so their plumes can be identified by elevated concentrations of SO_2 and NO_y (Ryerson et al., 1998; Ryerson et al., 2003). SO_2 enhancement can be a more reliable diagnostic of PPPs than NO_y since there are numerous sources of NO_x, but coal-fired power plants are dominant sources of SO_2 in eastern Texas (Ryerson et al., 2003; Neuman et al., 2009). Background SO_2 levels were consistently below 1 ppb, so this level of SO_2 is chosen as a threshold value for identifying PPPs.

In rural areas of northeastern Texas, power plants are also leading sources of CO and anthropogenic CO_2 (Nicks et al., 2003), even though CO is not strongly elevated in all PPPs. Airborne measurements in 2000 and 2006 showed that CO and CO_2 could be signatures of the

Martin Lake, Monticello, and Welsh plumes, the concentration enhancements of which completely overlap SO_2 and NO_y concentration enhancements at transects (Nicks et al., 2003). As the atmospheric lifetime of CO_2 is years, it is a conservative species in plumes. CO has a lifetime of one to two months in the atmosphere (Akimoto, 2003), thus serving as another conservative species in PPPs. CO emissions from Martin Lake, Monticello, and Big Brown, which were significantly underestimated in a previous emission inventory (1999), were more accurately estimated in TexAQS II (Peischl et al., 2010).

Of the 18 WP-3 flights during TexAQS II, the 16 September and 25 September flights measured successive cross-wind transects of PPPs from multiple power plants and the 19 September flight measured the Parish plume (Fig. 1 and Table S3). The 16 September flight (11:00 to 15:00 local time) observed transects of plumes from the Martin Lake, Pirkey, Monticello, and Welsh power plants at successive downwind distances (Fig. 2a). Since the three plumes transported northward through rural areas devoid of other large anthropogenic SO_2 sources, SO_2 concentration enhancements clearly denote plume locations (Fig. 2a). Pirkey is located several km north-northeast (downwind) of Martin Lake, so their plumes cannot be distinguished on this flight after the first Martin Lake transect; we refer to the plume as Martin Lake (Ma-1) for simplicity. On 19 September, the WP-3 measured five plume transects of Parish in the Houston-Galveston Brazoria (HGB) metropolitan region. On 25 September (13:00 to 16:00 local time), the WP-3 measured two plume transects of Big Brown and Limestone under northerly flow, and two plume transects of Parish (Fig. S5 and S6 in supplementary materials). All transects on the three days occurred at altitudes of 600-700 m, well within the planetary boundary layer height of approximately 1500 m determined from measured temperature profiles.

The exception was five transects (Ma-4 to Ma-8) of the Martin Lake plume at different heights but at the same downwind distance on 16 September.

3 Model setup and input parameters

Atmospheric chemistry for the episode was simulated by the Community Multiscale Air Quality (CMAQ) model (Byun and Schere, 2006) version 4.7 (Foley et al., 2010), using the CB05 chemical mechanism (Yarwood et al., 2005). Inline processing was applied to generate the meteorology dependent emissions properties (i.e., biogenic emissions) (Foley et al., 2010). After accounting for plume rise, most of the power plant emissions were modeled to be released between 200 and 600 m elevation (Fig. S7).

The model was configured with 34 vertical layers and three one-way nested domains. The outer two domains cover the continental U.S. (148×112 with 36 km grid resolution) and the eastern U.S (279×240 with 12km grid resolution) including all of Texas, respectively. The rectangular frame in Fig. 1 shows the fine domain with 4km grid resolution. A full description of the modeling configuration and performance for the 12km domain can be found in Appel et al. (2009). The CMAQ modeling for the 4km domain was from September 1-25, 2006, which covers the days with WP-3 plume measurements.

Meteorology for the episode was simulated by the Fifth-Generation NCAR/Penn State Mesoscale Model (MM5) (Grell et al., 1994) version 3.7.4 for the 36 km domain. For the inner domains (12km and 4km modeling domains), the Weather Research and Forecasting Model (WRF) version 3.0 (Skamarock et al., 2008) was used because of its lower biases in simulated wind and temperature than MM5. Both models had 34 vertical layers extending from the surface

up to 100 hPa. WRF was applied with Asymmetric Convective Model 2 PBL model (Pleim, 2007), Pleim-Xiu Land Surface Model (Xiu and Pleim, 2001), Dudhia shortwave radiation scheme (Dudhia, 1989), RRTM longwave radiation scheme (Mlawer et al., 1997), Kain-Fritsch 2 subgrid convective scheme (Kain, 2004), and the Thompson microphysics scheme (Thompson et al., 2004). MM5 used similar physical schemes. The consistency between MM5 and WRF for the modeling domains was tested and verified (Appel et al., 2009). Meteorological fields were converted to CMAQ-ready format by MCIP version 3.4.2 (Otte and Pleim, 2009).

Emission inputs for the three modeling domains were generated by the Sparse Matrix Operator Kernel Emissions (SMOKE) model (EPA, 2006) based on the National Emission Inventory for 2005. Mobile emissions were projected to 2006 and actual Continuous Emission Monitoring System (CEMS) data were used for point sources. BEIS3.12 (Environmental Protection Agency Biogenic Emissions Inventory System 3.12) (http://www.epa.gov/asmdnerl/biogen.html) was applied to compute the biogenic emissions.

NO_y species in the CB05 chemical mechanism are NO, NO_2, NO_3, N_2O_5, HONO, HNO_3, PNA (peroxynitric acid), PAN (peroxyacetyl nitrate), PANX (C3 and higher peroxyacyl nitrates), and NTR (organic nitrate). The sum of all these species (with N_2O_5*2) is the concentration of NO_y from the model.

The aqueous processing module in CMAQ (Walcek and Taylor, 1986) processes the absorption of gas-phase species and accumulation-mode aerosols separately. The gas-phase absorption into liquid water content of clouds depends on the thermodynamic equilibrium, whereas accumulation-mode aerosols are assumed to be completely activated to form cloud droplets. The dissociation of compounds into ions, oxidation of S(IV) to S(VI) by aqueous H_2O_2, O_3, Fe(III) and Mn(II) etc, and wet deposition are also processed in the model. For computational

efficiency, CMAQ does not transport cloud-aqueous concentrations separately from gas-phase concentrations between model grids. At the end of the cloud processing module, the cloud concentrations are removed and the mass of each species is passed to either gas-phase or aerosol concentrations.

In this study, the advection schemes used in processing pollutant transport by CMAQ are Piecewise Parabolic Method (PPM) (Colella and Woodward, 1984) and Yamartino-Blackman Cubic Scheme (YAM) (Yamartino, 1993). The Asymmetric Convective Model version 2 (ACM2) (Pleim, 2007) was used to simulate the vertical mixing of pollutants in CMAQ.

To identify and analyze the impact of each power plant, a zero-out simulation is run with the emissions of that facility removed from the base emission inventory. The difference between concentrations in the base simulation and each zero-out simulation represents the zero-out-contribution (ZOC) of that power plant.

4 Results and Discussion

During airborne measurement on the three days, ground temperature was 24.4-35.5°C (average: 29.0 °C) and surface wind was 0-7.2 m/s (average: 3.1m/s) at ground-based monitors in eastern Texas. At 600-700m above the ground level (WP-3 typical flying height), the observed ambient temperature was 23.7-30.3°C (average: 26.8 °C), wind speed was 1.6-12.0 m/s (average: 6.4 m/s) and no precipitation was observed. The height of the planetary boundary layer (PBL), determined from the vertical profiles of equivalent potential temperature for the three days, was about 1500 m on 16 September and about 1000 m in the HGB region on 19 and 25 September.

The CEMS-reported SO_2 and NO_x emissions of big power plants in the eastern U.S. were previously evaluated based on with WP-3 measurements of PPPs in 2004 (Frost et al., 2006). Since the emitted NO_x in PPPs can quickly be oxidized to NO_z ($NO_z = NO_y - NO_x$), the observed enhancements of NO_y and SO_2 serve as the basis for evaluation. The strong correlation between NO_y and SO_2 for all first plume transects (R^2=0.68~0.98) suggests that the power plants were the dominant sources of these gases there. The three ratios of these plants show strong consistency within the uncertainties of the measurements, although the model slightly under-predicts SO_2/NO_y ratios (Table 2). Likewise, previous studies have reported strong consistency between CEMS(SO_2/NO_x) and OBS(SO_2/NO_y) (Frost et al., 2006; Ryerson et al., 2003; Ryerson et al., 2001).

4.1 Evaluation of plume dispersion and transport

On 16 September, the WP-3 observed mostly southerly winds with average wind speeds of 6.9 m/s. The southerly winds allowed PPPs of Monticello and Welsh to remain distinct in both model and observation (Fig. 2) but caused the Martin Lake and Pirkey plumes to coincide since the two plants are just 18.5 km apart. Maximal SO_2 enhancements for each plume transect were used to identify the plume centers to enable comparative analyses of observations and modeling results. The plumes produced by CMAQ mostly have similar spatial extent to the measured plumes on 16 September (Fig. 2), 19 and 25 September (Fig. S5-S6). The wind speed and direction in the model were more homogeneous than observed winds, resulting in slight differences between modeled and observed locations of the plumes and plume centers (Fig. 2).

The high-resolution aircraft observations were compared with the model outputs extracted from the corresponding grid cells, adjusted to align the modeled and measured plume peak

locations as necessary. Since the aircraft was flying consistently at approximately 100 m/s at each plume transect, each gridline interval in Fig. 3 (40 seconds) is equal to the spatial distance of 4 km (one grid cell).

The 16 September flight path proceeded northward in 14 successive crosswind (east-west) transects, the first 12 of which intercepted the Martin Lake (and Pirkey) plumes (Ma-1 to Ma-12 in Fig. 2; Ma-4 to Ma-8 are increasing altitudes at the same transect) and the last four of which intercepted the Monticello and Welsh plumes (Fig. 2). The extensive observation of the Martin Lake plume provides a unique opportunity to examine plume evolution from the emission stack until dilution to background levels. Comparisons between modeled and observed SO_2, NO_y, and CO mixing ratios are shown for each successive plume transect in Fig. S1- S4.

At the first transect of the Martin Lake, Monticello, and Welsh plumes, the model generates lower peak SO_2 and NO_y concentrations and wider plumes than was observed. This likely reflects the inability of the 4-km resolution model to resolve subgrid-scale plume structure in the initial formation of a plume. No subgrid or Plume-in-Grid (PinG) was used in the modeling.

The model captured the observed extent of CO at each plume transect, slightly under-estimating the peak values (Fig. 3). The modeled SO_2 (18ppb) at Ma-2 matched the observed peak (23 ppb) closely as subgrid effect weakened and the plume width was larger than one grid cell. As the plume transported to Ma-3, the modeled SO_2 (14ppb) was higher than the observed peak (7ppb). The modeled SO_2 at plume center was consistently higher than the observed while the background SO_2 matched observations.

The measured CO at the plume center declined only from 240 (Ma-1) to 150 ppb (Ma-3). However, SO_2 was observed to decline by more than a factor of 10 from Ma-1 to Ma-3, indicating rapid loss.

Ma-4 through Ma-8 observed the Martin Lake plume at the same downwind distance (53 km) but flew at different altitudes (Ma-4 to Ma-8 of Fig. S1; Table S3). SO_2 emission from Martin Lake was modeled to occur mostly at 400 m, accounting for the stack height and plume rise (Fig. S7). At 1800 m (Ma-4), which was near the top of the PBL, no enhancement of SO_2, NO_y, or CO was simulated but a weak SO_2 plume was observed, implying that the model failed to capture some of the observed upward transport. At lower flight altitudes (between 660 and 300 m, corresponding to Ma-6 to Ma-8 in Fig. S1), the model effectively simulated plume extent. The comparisons between the modeled and observed SO_2, CO and NO_y species on 19 and 25 September are shown in Fig.S1 to S4.

4.2 Correlations between conservative and non-conservative species

In this section, we explore the correlation between conservative and non-conservative species from the observed plume concentrations. The correlations are presented by the slopes and R^2 of the least-square-fit between conservative and non-conservative species. At the time scale of PPP transport (a few hours), CO and CO_2 are expected to experience similar dispersion and minimal loss to chemistry or deposition, leading to near constant slopes of CO to CO_2. CO and CO_2 concentrations were strongly correlated within the Martin Lake and Monticello plumes and the slopes of CO to CO_2 held steady as both plumes aged (Fig. 4) (for Ma-1 to Ma-3 and Ma-6 to Ma-12, slopes of the least square fit: 0.58~0.71 ppb/ppm, R^2: 0.89~0.96; for Mo-1 to Mo-4, slopes of the least square fit: 4.3~5.3 ppb/ppm, R^2:0.77~0.94), indicating the same extent of dispersion of CO and CO_2. For the Welsh and Big Brown plumes, only the first one or two transects had a strong correlation between CO and CO_2, with the later transects likely affected by nearby CO or CO_2 emissions. Due to the strong interference from HGB urban emissions, no clear correlation between CO and CO_2 could be found in the Parish plume.

Concentrations of SO_2 and reactive nitrogen species in PPPs are strongly affected by chemical reactions, heterogeneous conversion, deposition, dispersion, and cloud processing. Dispersion is expected to have the same extent of impact on both conservative (e.g., CO_2 and CO) and non-conservative species (e.g., SO_2, NO_x, HNO_3, and PAN). Thus, the variations of slopes between non-conservative and conservative species reflect the impact of plume chemistry, deposition and heterogeneous processing on non-conservative species.

Given that SO_2 and CO_2 were observed to be strongly correlated in all plumes, CO_2 could serve well as a signature of PPPs. However, since CO_2 is not modeled by CMAQ, CO is selected as the conservative species for model-observation comparisons. CO is a signature emission of some but not all power plants in Texas. In observed PPPs, only in the Martin Lake and Monticello plumes could the strong correlations between the non-conservative species and CO be found at all transects. In Big Brown and Welsh plumes, SO_2 strongly correlates with CO, only in the initial transects.

4.3 Evaluation of SO_2 plume evolution

MODIS images (http://ladsweb.nascom.nasa.gov/browse_images/) and aircraft-observed photolysis rates indicate scattered cloudiness over eastern Texas on 16 September and clear skies on 19 and 25 September (Figs. S8 and S9). Relative humidity reached saturation between 1800~1000 m during descents through the Martin Lake plume on 16 September, implying clouds distributed at that altitude and potentially interacting with the plume (Ma-4 and Ma-5 in Fig. S1). The model successfully simulated the MODIS-observed distribution of scattered clouds over northeastern Texas on that day (Fig. S9 and S10), but placed them predominately between 2500

and 4000 m altitude (Fig. S10), well above the plumes (Fig. S11 and S12). Thus, no significant cloud processing was modeled to occur in the base modeling.

Under the clear skies of 19 September (Fig.S9), the normalized SO_2 to CO ratio from the model and the normalized SO_2 to CO_2 ratio from the observation matched closely for the Parish plume, showing slow SO_2 loss (Fig. 8). At the plume age of 11 hours, only 25% SO_2 was removed in both the modeling and the observation. SO_2 loss was also slow in the Big Brown and Parish plumes during cloud-free days (Fig. 5 and 8). Thus, the model can capture SO_2 evolution when no cloud processing occurs.

However, plume observations demonstrate rapid loss of SO_2 in the 16 September plumes (Martin Lake, Monticello and Welsh) where scattered clouding was observed (Fig. 5). For the Martin Lake transects, the decreasing trend of SO_2/CO fits to an exponential function with a first-order loss rate of 0.38 $hour^{-1}$, the inverse of which is a lifetime of 2.6 hours (R^2=0.94) (Fig. 5). SO_2/CO from the model decreases far slower as plume ages with a loss rate of 0.016 $hour^{-1}$ (lifetime of 62.5 hours), which suggests the model significantly underestimates SO_2 loss for the Martin Lake plume. Similarly, for the Monticello plume, the curve fit of observed SO_2/CO indicates an SO_2 lifetime of 2.7 hours compared to a modeled SO_2 lifetime of 17.2 hours. Although SO_2 and CO were not strongly correlated in observations of the other plumes, diminishing SO_2/CO_2 ratios indicate that rapid SO_2 loss also occurred in the Welsh plume.

The lifetime of SO_2 against gas-phase oxidation by OH is a few days to one week, and SO_2 lifetime against dry deposition approximates one day in the boundary layer. Thus, gas-phase oxidation and dry deposition are insufficient to explain the rapid loss of SO_2 in the 16 September plumes.

Could CMAQ have simulated the rapid SO_2 loss on 16 September if the meteorological model had placed the clouds at lower altitudes in contact with the PPPs? The cloud module in CMAQ includes two mechanisms for removing pollutants: aqueous chemical reactions and scavenging and wet deposition. SO_2 absorption into cloud droplets and subsequent oxidation are explicitly represented. The absorption is governed by thermodynamic equilibrium, followed by oxidation of aqueous S(IV) to S(VI) by H_2O_2, O_3, metal ions (Fe(III) and Mn(II)), and methylhydroperoxide (MHP), and peroxyacetic acid (PAA). Since no precipitation was observed or modeled during the airborne measurements, pollutants were not expected to be scavenged.

Cloud parameters of meteorological inputs are perturbed to diagnose how efficiently pollutants such as SO_2 and NO_y would have been removed from plumes. Specifically, the cloud bottom height in the meteorological field on 16 September is adjusted to 1000 meters so that the plumes interact with clouds during their transport. Liquid water content (QC) is the cloud parameter determining the extent of the pollutant aqueous processing must be larger than 0.01 g/kg to trigger the cloud aqueous module. In the perturbation cases, we uniformly increase QC to 0.05 g/kg ($\approx$0.05g/m^3, equivalent to fog), 0.5 g/kg ($\approx$0.5g/m^3, equivalent to stratocumulus clouds), and 5 g/kg (($\approx$5g/m^3, equivalent to cumulonimbus clouds).

In the base modeling for the Martin Lake plume, only 11% SO_2 is removed in the model (the normalized SO_2/CO decreased to 0.89 from Ma-1 to Ma-12). In the QC_0.05 case, 25% of SO_2 is removed during that span, far short of the observed 92% SO_2 removal (Fig. 5). The cloudier scenarios yield 66% (QC=0.5 g/kg) and 81% (QC=5.0 g/kg) SO_2 removal, still below the observed rate.

Five S(IV) oxidation reactions are explicitly implemented in the cloud aqueous module, i.e. H_2O_2, O_3, metal (Fe(III) and Mn(II)), MHP, and PAA oxidations. In QC_0.05, S(IV) oxidation

is dominated by H_2O_2 oxidation, with 96.2% of S(IV) oxidation occurring by H_2O_2 in the Martin Lake plume. Only about 1.7% of S(IV) oxidation was by the metal ions.

In the default CMAQ cloud module, Fe(III) and Mn(II) are uniformly set to 0.01 and 0.005 ug/m^3, representative of the background atmosphere. However, power plants are major emission sources of particulate metals (Alexander et al., 2009), and may have much higher levels of Fe(III) and Mn(II), thus potentially enhancing the aqueous oxidation of sulfur in PPPs. In another perturbation case, both Fe(III) and Mn(II) concentrations are increased by a factor of 10 in the QC_0.05 case (called QC_METAL hereafter). The increase of Fe(III) and Mn(II) is within the range of metal ion concentrations measured in fogs and cloud water (Raja et al., 2005;Parazols et al., 2006). SO_2 removal in QC_METAL was more rapid than that of QC_0.05 (Fig. 5). At the last plume transect, SO_2 decreased by 33%, compared to the 25% SO_2 removal in QC_0.05, suggesting the increased metals in plume lead to more rapid SO_2 oxidation. Thus, some combination of enhancements in cloud liquid water content and metals concentrations may help explain the observed rapid SO_2 loss rates in the cloudy day plumes.

Few studies have observed rapid SO_2 loss in anthropogenic plumes, though similar rates of SO_2 loss have been found in volcanic plumes (Oppenheimer et al., 2010;Rodríguez et al., 2008). These studies proposed that cloud aqueous processing is the mechanism for the rapid SO_2 removal. The comprehensive airborne measurement of plume concentrations and meteorological parameters supported by satellite images in this study confirms that the cloud processing caused the rapid SO_2 loss. SO_2 uptake by cloud droplet and subsequent aqueous oxidation are a major challenge to models. Earlier studies have also found that models can underestimate SO_2 loss rates in clouds (Crutzen and Lawrence, 2000;Kreidenweis et al., 1997).

4.4 Evaluation of plume chemistry of reactive nitrogen

In PPPs, HNO_3, NO_3, N_2O_5, PAN, and other organic nitrates are formed via NO_x chemical reactions. Freshly emitted NO_x titrates O_3 and consumes OH, resulting in slow formation of HNO_3 and no formation of PAN in the initial plume (Karamchandani et al., 1998). As a plume dilutes, OH levels recover and HNO_3 and other products form from NO_x oxidation. Previous daytime observations of PPPs concluded that HNO_3 and PAN were the major (more than 90%) products of NO_x oxidation in PPPs (Neuman et al., 2006; Neuman et al., 2004; Ryerson et al., 2003; Ryerson et al., 2001). The observational data in this study also show that HNO_3 and PAN were the only two major oxidation products in PPPs, with NO_3 and N_2O_5 and other organic nitrates at least one order of magnitude lower in plume transects.

The measured and modeled NO_x, HNO_3, and PAN are shown for comparison in Fig. S2 and S3. NO_x was higher than HNO_3 until the plume transported 2.0 hours at Ma-7 and Ma-8. The model generally captured the observed evolution of reactive nitrogen species NO_x, HNO_3, and PAN in the plume, simulating the transition from NO_x to HNO_3 dominance and approximately matching the observed PAN levels. However, the simulated HNO_3 concentrations were higher than observed, implying over-prediction of HNO_3 formation or under-prediction of HNO_3 loss.

The oxidation of NO_x by radicals approximates as a first-order reaction if radical concentrations are assumed to be constant in the plume. The observed NO_x/CO fits to an exponential decay function (for Martin Lake, R^2=0.85; for Monticello, R^2=0.86; Fig. 6), corresponding to NO_x lifetimes of 2.6 and 1.2 hours for the Martin Lake and Monticello plumes, respectively. The NO_x lifetimes computed here are consistent with the NO_x lifetimes (1.0~1.6 hours) estimated for both plants in TexAQS 2000 (Neuman et al., 2004). The declining trends of NO_x/CO from the model and the observation closely match in the Martin Lake and Monticello

plumes, with discrepancies only in the initial transects due to the inability of the model to resolve subgrid-scale plume structure (Fig. 6).

The ratios of HNO_3/CO and PAN/CO are compared between the model and observations to explore chemical evolution in the Martin Lake and Monticello plumes. We find that the model captures the PAN formation very well, closely matching observed trends as the plumes age (Fig. 6 and Fig. S2). The modeled HNO_3/CO, however, was 0.7~6.6 times larger than observed. Given the good agreement between the modeled and observed NO_x oxidation and PAN formation, the HNO_3 gap between the model and the observation on the cloudy day implies that HNO_3, while formed during plume transport, was rapidly removed from the atmosphere, which is not captured by the model.

Unexpectedly rapid loss of NO_y has also been reported by some measurement studies of biomass burning (Takegawa et al., 2003) and PPPs (Neuman et al., 2004), but not in others (Ryerson et al., 2003). When NO_x is oxidized to other reactive nitrogen species, the reactive nitrogen may be removed from the atmosphere via rain scavenging, dry deposition, heterogeneous conversion to aerosol, and cloud processing, resulting in the loss of NO_y. Assuming a first-order decline of NO_y/CO (Fig. 7), the observed NO_y loss rate was 0.15 $hour^{-1}$ for the Martin Lake plume whereas the modeled NO_y loss rate was lower by a factor of 6 (0.026 $hour^{-1}$). For the Monticello plume, the observed NO_y loss rate (0.24 $hour^{-1}$) was 2.3 times the modeled. The observed NO_y/CO_2 in Martin Lake, Monticello, and Welsh plumes had a similar extent of NO_y loss, especially during the early plume age ($\leq$2 hours) when NO_y/CO declined by 40~50% (Fig. 7).

On 19 September, a cloud-free day, the model effectively simulates the observed slow removal of NO_y (Fig 8). NO_y loss on the cloudy day likely reflects deposition of highly soluble

HNO_3, since the other main NO_y constituents (NO_x and PAN) have low water solubility, cannot directly convert to aerosol, and have negligible dry deposition in plume. NO_x oxidation and thermal decomposition of PAN do not shift the gas-phase NO_y budget since their products are also gas-phase NO_y constituents. The measured NO_3^- was minor in the inorganic aerosol composition, indicating that the loss of HNO_3 to aerosol-NO_3^- was negligible under the high ambient temperatures (the measured average temperature was 28.9°C) and the lack of ammonia enhancement beyond levels needed to neutralize the sulfate in the PPPs (Nowak et al., 2010). Given that no wet precipitation was reported on the flight days, no rain scavenging is expected to have occurred. HNO_3 may have rapidly dissolved in cloud droplets if the plume interacted with a cloud, as is possible under the 16 September scattered cloudiness conditions discussed earlier.

In contrast to the SO_2 results, the cloud perturbation scenarios did not significantly impact modeled concentrations of NO_y species. Among NO_y species, HNO_3 is the only one to be processed by the cloud module. Even raising QC to 5.0 g/kg, there is no scavenging removal of HNO_3 since no wet deposition happens in the absence of precipitation. At the end of the cloud module, CMAQ returns the aqueous concentration of HNO_3 in cloud to either gas-phase or aerosol species, keeping the ratio of HNO_3 to total NO_y constant. This makes HNO_3 concentrations insensitive to QC in CMAQ.

4.5 Evaluation of O_3 simulation in PPPs

Various numerical models have been applied to simulate the O_3 chemistry of PPPs (Sillman, 2000;Springston et al., 2005;Frost et al., 2006;Zaveri et al., 2010). The simulation of the chemistry and transport of PPPs by 3-D models are widely used for assessing the effectiveness of emission controls of power plant pollutants. However, the performance of these models has most

often been evaluated with ground concentrations (Mauzerall et al., 2005;Vijayaraghavan et al., 2009;Godowitch et al., 2008a;Godowitch et al., 2008b).

The model overestimated background O_3 by 8~15 ppb during the flights (Table S4). Sensitivity modeling shows that boundary conditions were the biggest contributor to background O_3 levels. Thus, we focus on the differences (ΔO_3) between plume and background O_3 mixing ratios to assess model performance for O_3 formation from power plant plumes (Table S4).

The model accurately simulates that the Monticello and Welsh plumes shift from being depleted to being enriched in O_3 between transect 1 and 2, and predicts the transition to occur one transect sooner than observed for Martin Lake. All of these plumes traversed rural regions of northeastern Texas where biogenic isoprene is abundant. However, the model underestimates the amount of O_3 enrichment downwind by 20-70% (Ma-9 to Ma-12, Mo-2 to Mo-4, We-2 to We-4). The model also underestimates titration in the initial transects, reflecting the more rapid dilution of NO_x in the model.

Ozone Production Efficiency (OPE) quantifies the number of O_3 molecules formed per molecule of NO_x irreversibly oxidized to NO_z species (Liu et al., 1987). Here, we determine OPE in the model from the ratio ZOC_{O3} to ZOC_{NOz}, and in the measurements from the least square slope of O_3 versus NO_y-NO_x (NO_z) (Trainer et al., 1993;Kleinman et al., 2002;Griffin et al., 2004;Ryerson et al., 2003).

Observations of Martin Lake, Monticello, and Welsh transects show OPE steadily increasing, while O_3 production evolves from depletion to net formation, consistent with OPE trends from PPPs in the southeast U.S. (Ryerson et al., 2001) and in Texas in 2000 (Springston et al., 2005;Ryerson et al., 2003). OPEs from transects of these plumes (Ma-6, Mo-4, and We-4) at similar plume ages are compared in Fig. 9. OPEs for Monticello and Welsh were remarkably

similar (Fig. 9), reflecting approximately equal O_3 formation potentials of these facilities with similar NO_x emission rates (Table 1). Martin Lake emitted about two times as much NO_x as Monticello and Welsh, and thus exhibited a smaller OPE. Comparing younger transects, OPE in the Big Brown plume (Bi-1) was 1.7 at a plume age of 1.3 hours, lower than the similar-plume-age OPE of Martin Lake (2.6, Ma-3) and Welsh (4.6, We-2), but close to the OPE of Monticello (1.4, Mo-2). OPE could not be quantified in the subsequent Big Brown transect due to lack of correlation between O_3 and NO_z. The Parish plume exhibited an OPE of 4.4 at a plume age of just 0.6 hours, suggesting rapid O_3 formation under the influence of Houston region anthropogenic VOCs.

The model replicated the observed temporal trends in OPE but simulated maximum OPEs about a factor of 2 lower than observed (Table S4). Since the calculation of OPE implicitly assumes that NO_y is conservative in plumes, the rapid loss of NO_y observed but not modeled on the cloudy day (Table S4) may undermine the accuracy of the observation-based OPE.

It should also be noted that the model tended to under-predict measured isoprene concentrations. Observed isoprene concentrations averaged over all transects are higher than the modeled average by 51.3%. We perturb domain-wide isoprene emission rates by this factor in the model to investigate how much impact the isoprene discrepancy has on the O_3 formation in plumes. After perturbation, the simulated ZOC_{O3} has a maximum increase of 3 ppb (Ma-2). The maximum O_3 increase for Monticello and Welsh is 1.5 and 2 ppb, respectively. OPEs of Martin Lake, Monticello, and Welsh plumes would increase to 7.0, 7.8, and 5.9, respectively, closing roughly half of the gap between modeled and observed OPEs. For Big Brown, the OPE would increase by a factor of 1.4, and for Parish, the OPE would increase to 6 (Pa-2), exceeding the observed OPE (4.4).

5 Summary and Conclusions

A regional 3D photochemical model was applied with fine-grid resolution to simulate PPPs during three days of airborne measurement by NOAA's WP-3 aircraft in TexAQS II. In this modeling study, the modeled and airborne observed concentrations are compared in detail at each plume transect. Under steady wind meteorological conditions, the fine-scale (4km) CMAQ accurately simulated the transport and dispersion of PPPs despite lacking a plume-in-grid module.

SO_2 and NO_x show strong consistencies among the CEMS-reported emission data. In the Martin Lake and Monticello plumes, CO was strongly correlated with SO_2 and NO_y and could serve as a conservative tracer species to track plume evolution; CO_2 was strongly correlated with SO_2 and NO_y in all plumes but was not modeled by CMAQ. The trend in the least square slopes of pollutants relative to CO (CO_2) was used to assess species lifetime.

On clear-sky days (19 and 25 September), SO_2 and NO_y experienced slow evolution (loss) in the Parish and Big Brown plumes. Both the model and the observation were closely correlated in the ratios of SO_2 and NO_y to conservative species, suggesting the model well captured SO_2 and NO_y evolution in the plumes.

SO_2 was observed to be rapidly lost in the Martin Lake, Monticello, and Welsh plumes under scattered cloudiness on 16 September. The observation-based SO_2 lifetime was 2.6 and 2.7 hours for the Martin Lake and Monticello plumes, respectively. The detailed examination of the photolysis rate and relative humidity data suggested cloud-processing of PPPs caused the rapid SO_2 loss on 16 September. The original simulation did not show the apparent SO_2 loss since

PPPs resided below clouds in the model. Perturbing the cloud bottom heights to interact with the PPPs yielded modest rates of SO_2 removal via aqueous processing in the CMAQ cloud module. SO_2 removal in the model was still slower than the observed rapid loss, even after increasing cloud liquid water content and metals concentrations in cloud droplets to enhance SO_2 oxidation.

The simulation closely matched the observed NO_x oxidation rates. The observed NO_x lifetime for Martin Lake and Monticello plumes was 2.6 hours and 1.2 hours, respectively. The modeled PAN formation reflected the observed trend of PAN formation, while the modeled HNO_3 was 0.7~6.6 times higher than observed due to the rapid HNO_3 loss in observation on cloudy days. For the Martin Lake plume, the loss rate of NO_y has been quantified to be 0.148 $hour^{-1}$ in observation, faster than the modeled NO_y (0.026 $hour^{-1}$) by a factor of 6 since modeled NO_y loss was insensitive to aqueous processing in the absence of precipitation.

The model effectively simulated the transition between ozone titration and formation but tended to predict lower OPE than indicated by observations. The discrepancies of OPEs between the model and the observations could be explained by the observed rapid NO_y loss that biases high the observation-based OPE estimates.

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Table 1. Major power plants in eastern Texas

Facility	NO_x emission rate [a] (tons/hour)	SO_2 emission rate [a] (tons/hour)	Stack Height (m)
Martin Lake	2.02	10.37	138
Monticello	1.34	5.49	128
Welsh	0.95	2.21	172
Pirkey	0.58	0.21	160
Big Brown	0.84	13.09	122
Parish	0.33	2.74	183
Limestone	0.79	0.63	137

[a] emission rate is the hourly averaged CEMS data for Martin Lake, Monticello, and Welsh on 16 September, 2006, and for Parish, Big Brown, and Limestone on 25 September, 2006.

Table 2. CEMS-reported $E(SO_2)/E(NO_x)$ emission molar ratio, the observed SO_2/NO_y and the modeled ZOC_{SO2}/ZOC_{NOy} at the location of the first plume transect.

Plant	CEMS SO_2/NO_x	OBS SO_2/NO_y [a]	MODEL ZOC_{SO2}/ZOC_{NOy} [b]	Plume age (hours)
Martin Lake	3.05	3.94 (0.98)	3.30	0.7
Monticello	2.04	3.00 (0.86)	1.84	0.3
Welsh	1.10	1.20 (0.86)	1.08	0.4
Big Brown	8.94	10.95 (0.97)	9.73	1.3
Parish	5.28	6.83 (0.68)	5.18	0.6

[a] the values in brackets are the R^2 of least square fit of SO_2 versus NO_y

[b] $ZOC_{SO2} = SO_{2\ model,\ base} - SO_{2\ model,\ zero\text{-}out\ that\ plant}$, $ZOC_{NOy} = NO_{y\ model,\ base} - NO_{y\ model,\ zero\text{-}out\ that\ plant}$

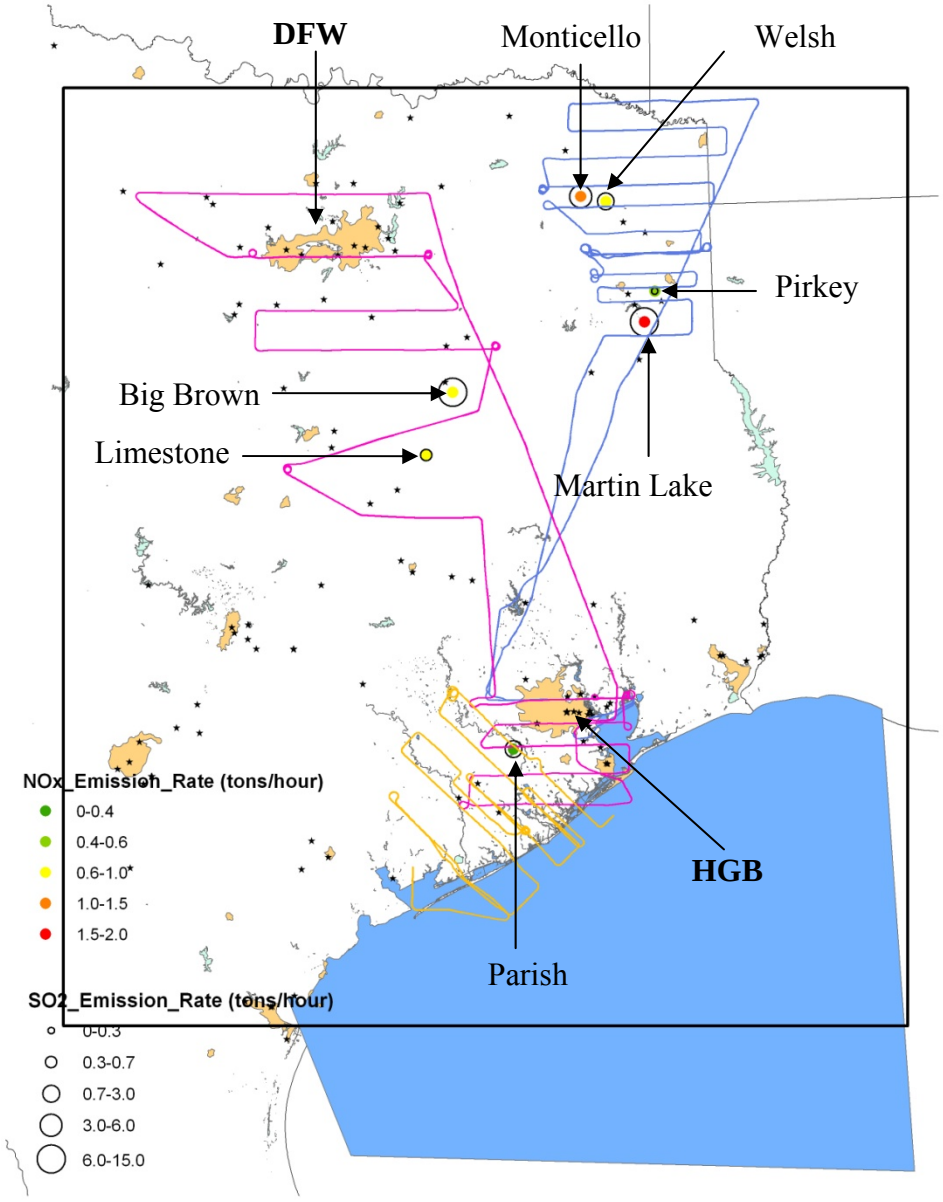


Fig. 1. WP-3 flight tracks (16 September in blue, 19 September in yellow, and 25 September in pink,) and power plants in eastern Texas. NO_x emission rates are shown by colors and SO_2 emission rates are indicated by size of circles. Rectangular frame shows the 4km modeling domain. Black stars are all other point sources in Texas. The Houston-Galveston-Brazoria (HGB) and Dallas-Forth-Worth (DFW) metropolitan areas are also shown.

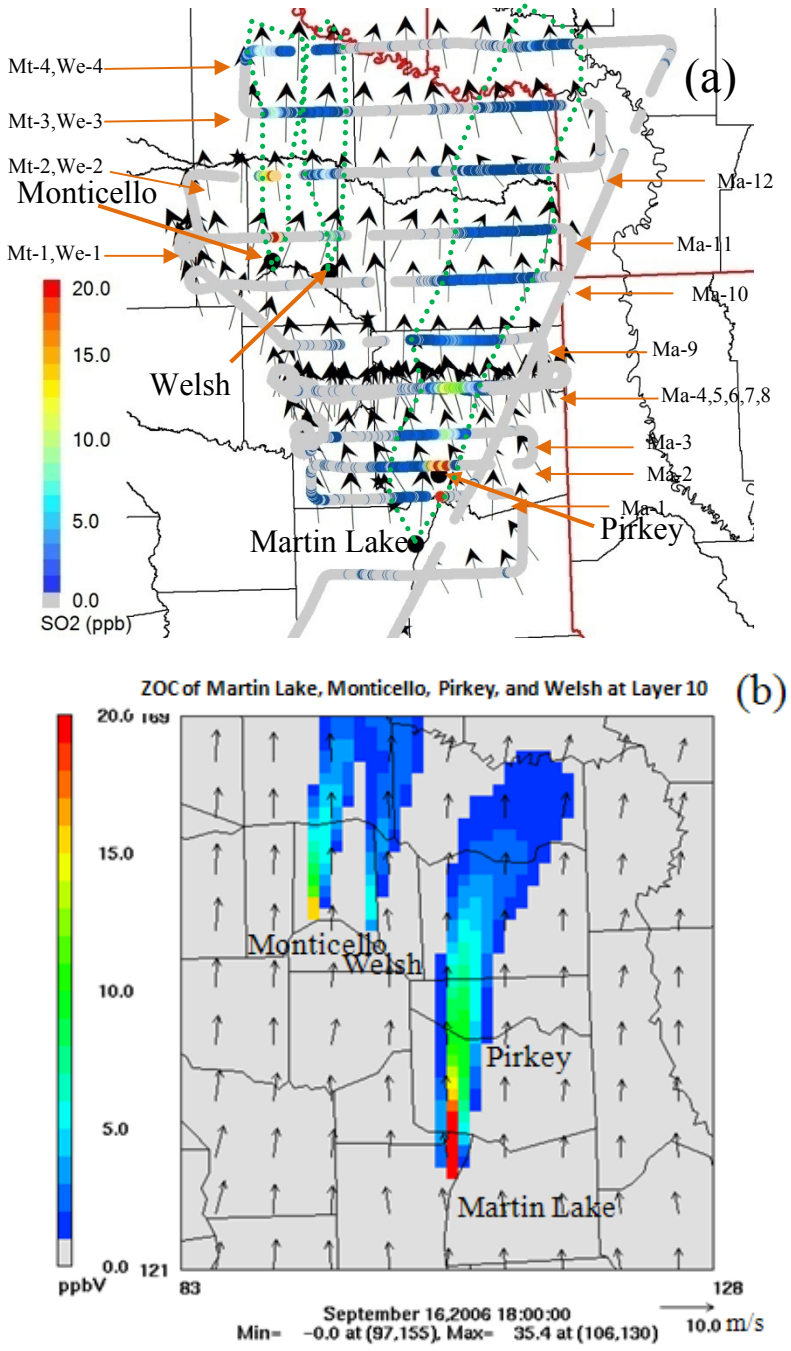


Fig. 2. (a) Observed PPPs of Martin Lake, Monticello, Pirkey, and Welsh on 16 September, 2006. The black dots show the locations of the power plants. PPPs are identified by measured SO_2 enhancement (color gradient in the figure), as outlined by the green dash lines. Measured wind vectors are presented on the plume transect. The geographical redlines are the state boarder-lines. (b) Simulated PPPs of Martin Lake, Monticello, Pirkey, and Welsh at 18:00 GMT (600~700 m) (local time: 12:00)

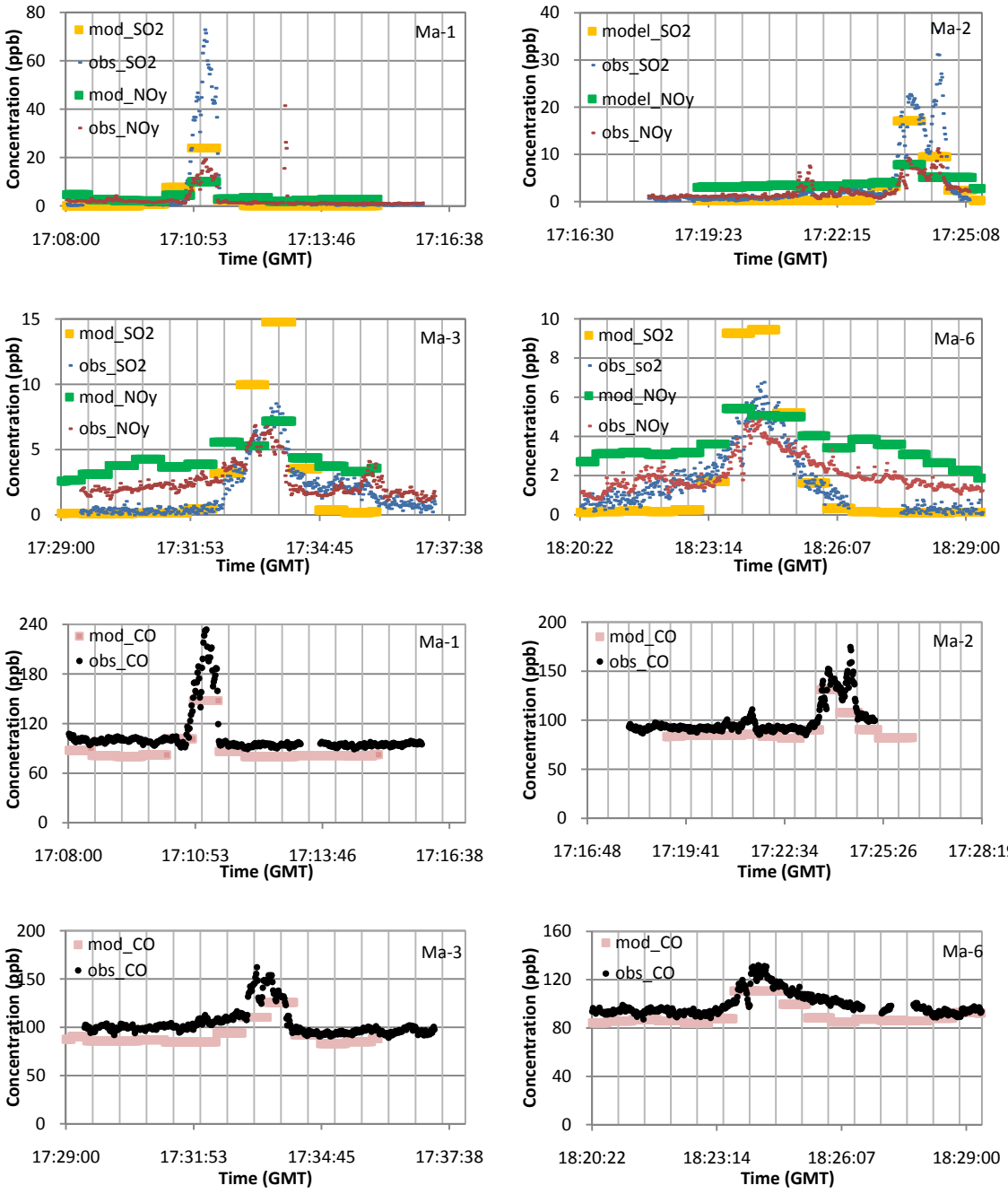
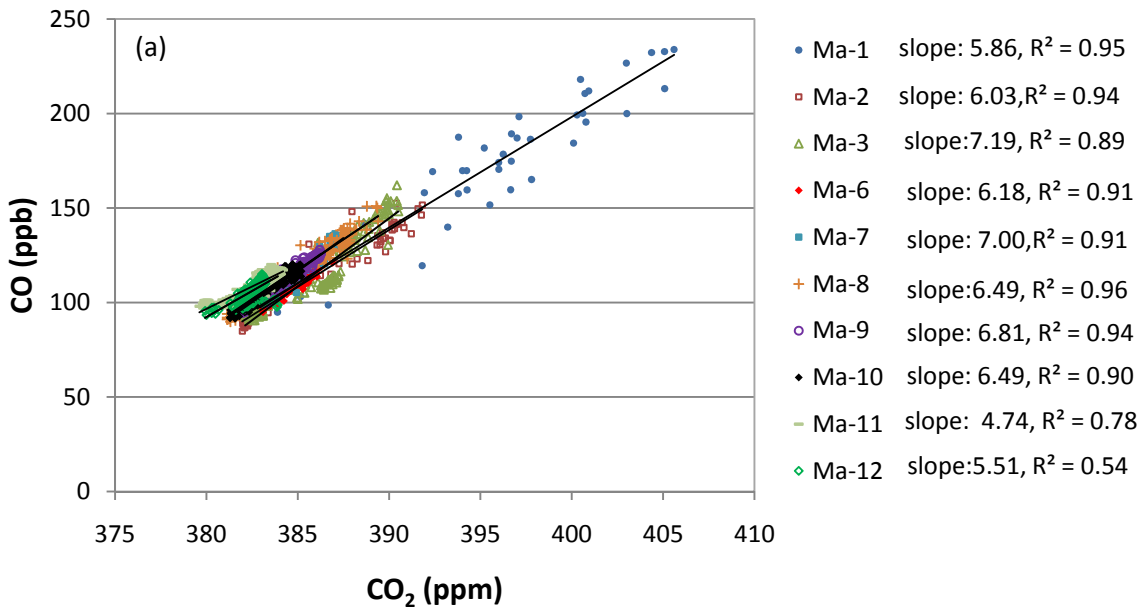


Fig. 3. The comparison of the modeled and observed SO_2, NO_y, and CO at plume transect of Ma-1, Ma-2, Ma-3 and Ma-6. The modeled SO_2, NO_y, and CO are labeled as yellow, green, and purple flat lines, respectively. The observed SO_2, NO_y, and CO are labeled as blue, red and black dots. The Horizontal coordinate is time scale in GMT (local time = GMT - 6 hours) and vertical coordinate is concentration (ppb). Transect names listed in Table S3 of the manuscript are labled in each subplot. Fig. S1-S4 summarize the comparisons for all the plume transects



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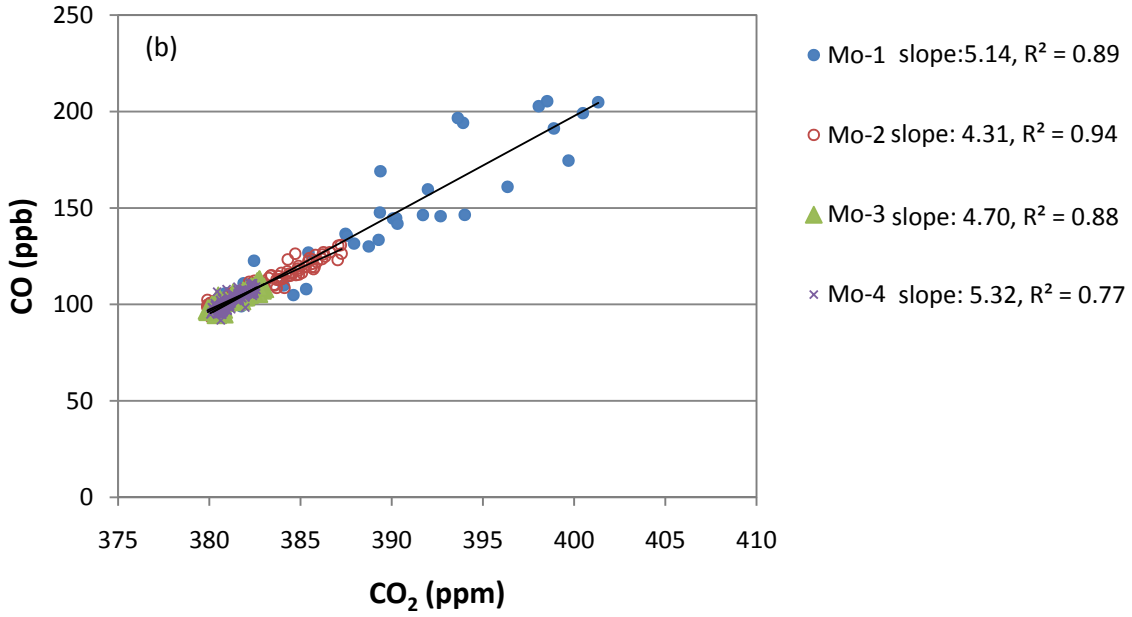

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3 Fig. 4. Scatter plot of CO (ppb) versus CO_2 (ppm) from plume transects (a) Martin Lake (Ma-1
4 to Ma-12), and (b) Monticello (Mo-1 to Mo-4). The unit of the slopes from the least square fits is
5 ppb/ppm.

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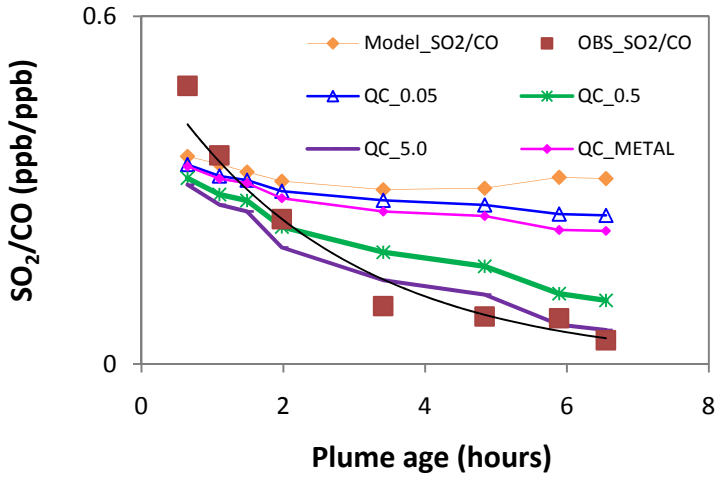


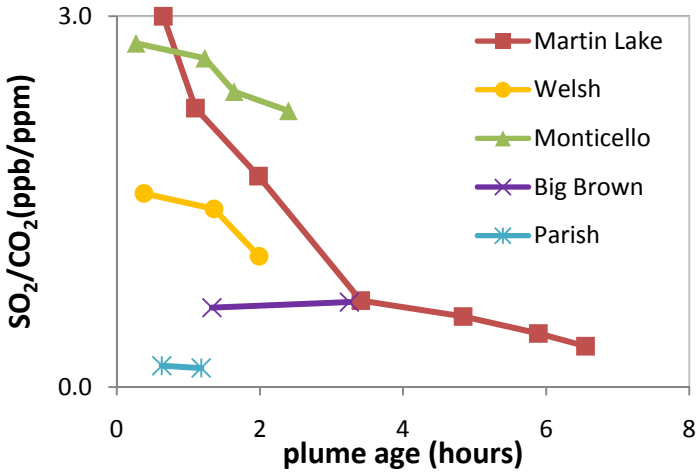


Fig. 5. (a) The observed least square slopes of SO_2 to CO (red square) and modeled ZOC_{SO2}/ZOC_{CO} (blue diamond for the base case, green dot for the adjusted cloud case) as a function of plume age at each transect of the Martin Lake plume (16 September). The observed SO_2 loss rate was 0.38 $hour^{-1}$ (R^2=0.94, SO_2 lifetime: 2.6 hours); the modeled SO_2 loss rate was 0.016 $hour^{-1}$ (R^2=0.36, SO_2 lifetime: 62.5 hours). The SO_2 to CO slopes for each perturbation case are also plotted accordingly. (b) The observed least square slopes of SO_2 to CO_2 (ppb/ppm) for the five plumes; Martin Lake, Welsh, and Monticello plumes were observed on September 16 (cloudy day), Big Brown and Parish plumes were made on 25 September (sunny day).

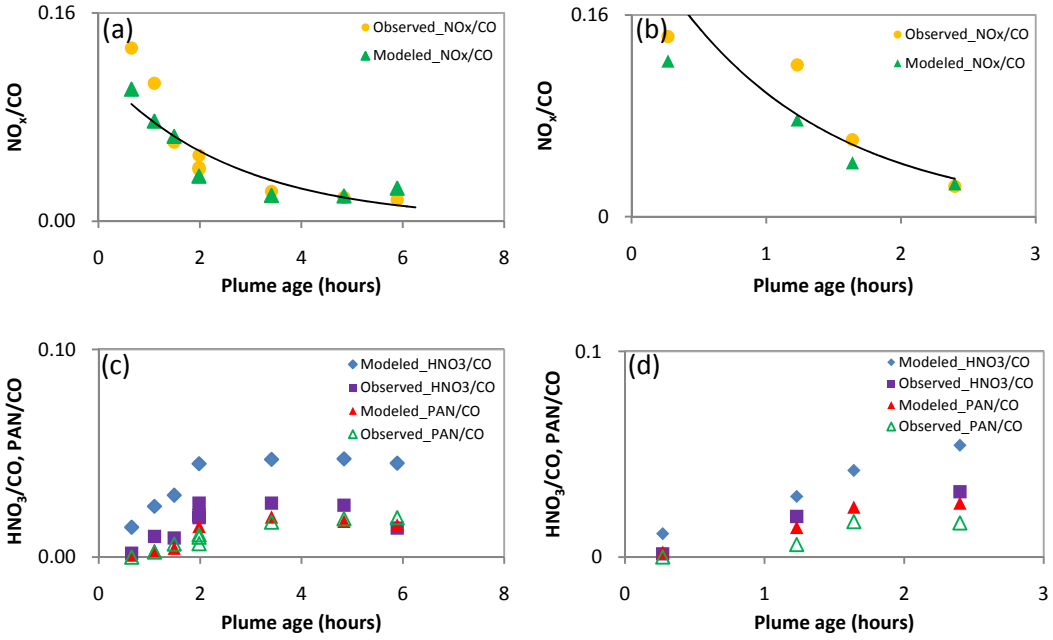


Fig. 6. Observed (yellow circle) and modeled (green triangle) NO_x/CO (a) for the Martin Lake plume, (b) for the Monticello plume. The observed NO_x oxidation rate was 0.38 hour^{-1} (R^2=0.85) for the Martin Lake plume and 0.84 hour^{-1} (R^2=0.86) for the Monticello plume. Observed and modeled PAN/CO and HNO_3/CO, (c) for the Martin Lake plume, (d) for the Monticello plume.

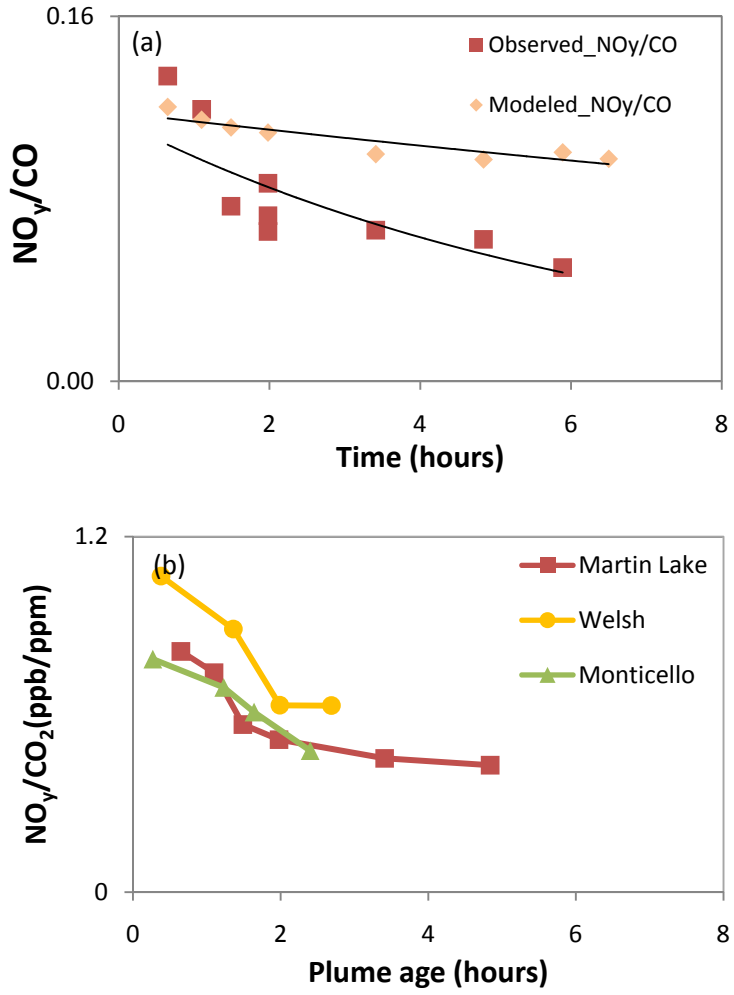
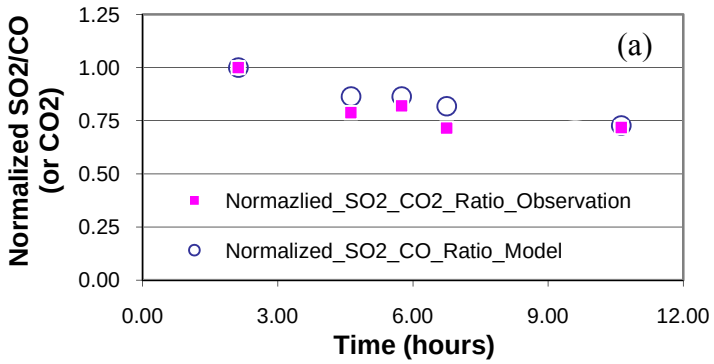


Fig. 7. (a) The observed least square slopes of NO_y to CO (red square) and modeled ZOC_{NOy}/ZOC_{CO} (orange diamond for the base case, green dot for the adjusted cloud case) as a function of plume age at each transect of the Martin Lake plume; the observed NO_y loss rate was 0.145 $hour^{-1}$ (R^2=0.69) and the modeled NO_y loss rate was 0.026 $hour^{-1}$ (R^2=0.48). (b) The observed least square slopes of NO_y to CO_2 (ppb/ppm) for the Martin Lake, Monticello, and Welsh. NO_y in the least-square fits was directly measured and not the sum of measured reactive nitrogen species.



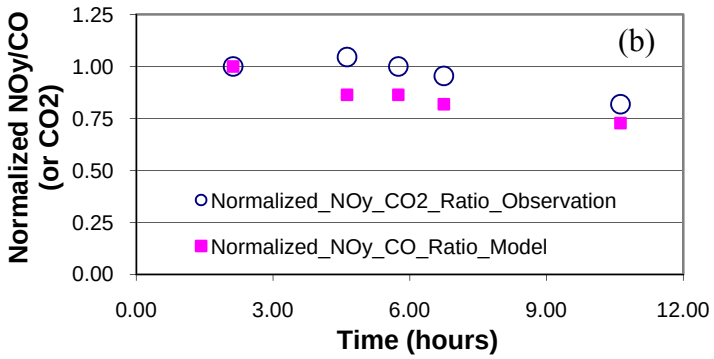


Fig. 8. The least-square-fit slopes of SO_2 to CO_2 (from the observation) and SO_2 to CO (from the model) (a), and the least-square-fit slopes of NO_y to CO_2 (from the observation) and NO_y to CO (from the observation) (b). All slopes are normalized to the slope at the first transect.

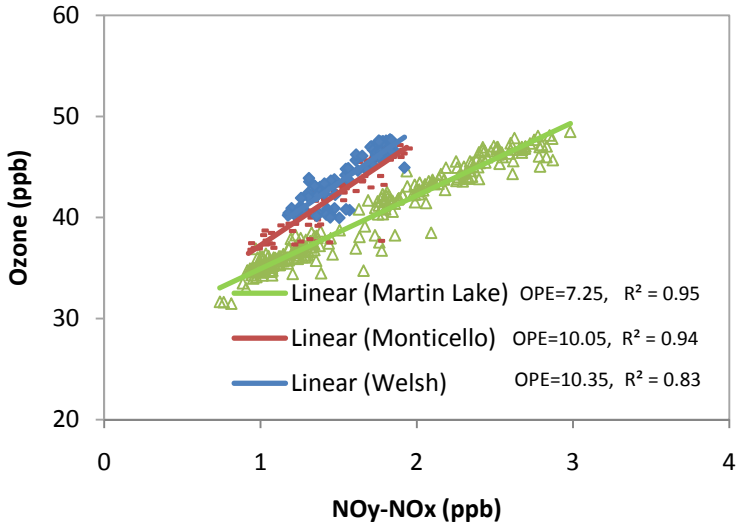


Fig. 9. O_3 versus NO_y-NO_x from the transects of Martin Lake (Ma-6, plume age of 2.0 hours), Monticello (Mo-4, plume age of 2.4 hours), and Welsh (We-4, plume age of 2.7 hours). The slopes from the least square fits indicate the observation-based estimates of OPE from each plume transect.